letters to nature

- 12. Moissette, A., Dubessy, L. Cathlineau, M., Ramanaidou, F. & Le Gleuher, M. Fluids with high acidity evidenced by Raman analysis of sulfate and bisulfate ions in fluid inclusions from a gold showing (Brazil). IAGOD Symp. Progm. Abstr. A18 (1990).
- 13. Rosasco, G. J. & Roedder, E. Application of a new Raman microprobe spectrometer to nondestructive analysis of sulfate and other ions in individual phases in fluid inclusions in minerals. Geochim Cosmochim. Acta 43, 1907-1915 (1979).
- 14. Dubessy, J., Boiron, M., Moissette, A., Monnin, C. & Stretenskaya, N. Determinations of water, hydrates, and pH of fluid inclusions by micro-Raman spectrometry. Eur. J. Min. 4, 885-894 (1992). 15. Plummer, L. N., Parkhurst, D. L., Fleming, G. W. & Dunkle, S. A. A Computer Program Incorporating
- Pitzer's Equations for Calculation of Geochemical Reactions in Brines (Water Res. Invest. Rep. 88-4153, US Geol. Surv., Reston, VA, 1988).
- 16. Turner, D. J. Raman spectral study of bisulphate ion hydration. J. Chem. Soc. Faraday Trans. 68, 643-648 (1972)
- 17. Rudolph, W. Structure and dissociation of the hydrogen sulphate ion in aqueous solution over a broad temperature range: a Raman study. Z. Phys. Chem. 194, 73-95 (1996).
- 18. Benison, K. C. Surface water paleotemperatures from Permian Nippewalla Group halite, Kansas Carbonates Evaporites 10, 242-248 (1995).
- 19. Benison, K. C. & Goldstein, R. H. Permian paleoclimate data from fluid inclusions in halite. Chem Geol. (in the press).
- 20. Hall, D. L., Bodnar, R. J. & Craig, J. R. Evidence for postentrapment diffusion of hydrogen into peak metamorphic fluid inclusions from the massive sulfide deposits at Ducktown, Tennessee. Am. Mineral. 76, 1344-1355 (1991).
- 21. Marvogenes, J. A. & Bodnar, R. J. Hydrogen movement into and out of fluid inclusions in quartz:
- Experimental evidence and geologic implications. Geochim. Cosmochim. Acta 58, 141-148 (1994). 22. Goldstein, R. H. & Reynolds, T. J. Systematics of Fluid Inclusions in Diagenetic Minerals 199 (Short Course 31, SEPM (Soc. for Sed. Geol.), Tulsa, OK, 1994).
- 23. Delmelle, P. & Bernard, A. Geochemistry, mineralogy, and chemical modeling of the acid crater lake of Kawak Ijen Volcano, Indonesia. Geochim. Cosmochim. Acta 58, 2445-2460 (1994).
- 24. Rowe, G. L. The Acid Crater Lake System of Poas Volcano, Costa Rica: Geochemistry, Hydrology, and Physical Characteristics. Thesis, Pennsylvania State Univ. (1991).
- 25. Nordstrom, D. K., Valentine, S. D., Ball, J. D., Plummer, L. N. & Jones, B. F. Measurement of negative pH values and high metal concentrations in extremely acidic mine waters from Iron Mountain, California. Geol. Soc. Am. Abstr. Progm. 23, 383 (1991).
- 26. Nordstrom, D. K. in Acid Sulfate Weathering (eds Kittrick, J. A., Fanning, D. S. & Hossner, L. R.) 37-56 (Spec. Publ. 10, Soil Sci. Soc. Am., Madison, WI, 1982).
- 27. Berendsen, P. Permian red bed copper mineralization in south-central Kansas. Acad. Sci. Trans. Abstr 83, 122 (1980).
- 28. Hovorka, S. Depositional environments of marine-dominated bedded halite, Permian San Andree Formation, Texas. Sedimentology 34, 1029-1054 (1987).
- 29. Lowenstein, T. K. Origin of depositional cycles in a Permian "saline giant": the Salado (McNutt zone) evaporites of New Mexico. Geol. Soc. Am. Bull. 100, 592-608 (1988).
- 30. Presley, M. W. Evolution of Permian evaporite basin in Texas panhandle. Bull. Am. Assoc. Petrol. Geol. Bull. 71, 167-190 (1987).

Acknowledgements. We thank T. K. Lowenstein and R. J. Bodnar for suggestions that improved this manuscript. This work was partially funded by NSF, GSA and Sigma Xi.

Correspondence and requests for materials should be addressed to K.C.B. (e-mail: kathleen.benison@ cmich.edu)

Boreal forest plants take up organic nitrogen

Torgny Näsholm, Alf Ekblad*, Annika Nordin, Reiner Giesler*, Mona Högberg* & Peter Högberg*

Department of Forest Genetics and Plant Physiology and * Section of Soil Science, Department of Forest Ecology, Swedish University of Agricultural Sciences, S-901 83 Umeå, Sweden

Plant growth in the boreal forest, the largest terrestrial biome, is generally limited by the availability of nitrogen. The presumed cause of this limitation is slow mineralization of soil organic nitrogen^{1,2}. Here we demonstrate, to our knowledge for the first time, the uptake of organic nitrogen in the field by the trees Pinus sylvestris and Picea abies, the dwarf shrub Vaccinium myrtillus and the grass Deschampsia flexuosa. These results show that these plants, irrespective of their different types of root-fungal associations (mycorrhiza), bypass nitrogen mineralization. A trace of the amino acid glycine, labelled with the stable isotopes ¹³C and ¹⁵N, was injected into the organic (mor) layer of an old successional boreal coniferous forest. Ratios of ¹³C:¹⁵N in the roots showed that at least 91, 64 and 42% of the nitrogen from the absorbed glycine was taken up in intact glycine by the dwarf shrub, the grass and the trees, respectively. Rates of glycine uptake were similar to those of ¹⁵N-ammonium. Our data indicate that organic nitrogen is important for these different plants, even when they are competing with each other and with non-symbiotic microorganisms. This has major implications for our understanding of the

effects of nitrogen deposition, global warming and intensified forestry.

Current conceptual models of nitrogen cycling in boreal forests are based on the assumption that mineralization of organic nitrogen is a prerequisite for plant nitrogen acquisition. Laboratory studies have shown, however, that some plants can use amino acids^{3,4} and proteins⁵, thereby bypassing the common mineralization pathway. This has not been demonstrated in the field, where plants compete with soil microorganisms held to be superior in using organic substrates. It has been proposed that plants forming ericoid mycorrhiza and ectomycorrhiza are especially efficient at using organic nitrogen⁶. This would give them a competitive advantage over plants forming arbuscular mycorrhiza and over non-mycorrhizal plants, and help explain their dominance in boreal forests⁶, in which levels of organic nitrogen are typically high in the soil (compare with Table 1). If so, anthropogenic inorganic nitrogen deposition, climate warming and intensified silviculture could alter the composition of plant communities by promoting a shift from organic- towards inorganic-nitrogen nutrition. Again, field evidence demonstrating that organic-nitrogen uptake occurs and is exclusive to some groups of plants is lacking. The interaction between the abiotic soil system, mycorrhizal fungi, other microorganisms and plants is complex⁷, and studies showing higher gross turnover rates of inorganic nitrogen pools than previously assumed⁸ call for intensified studies of the nitrogen cycle, taking into account higher levels of complexity⁹.

The lack of data showing organic nitrogen uptake by plants in the field relates to methodological problems. Classic ¹⁵N-tracer technique cannot separate uptake of intact amino acids from uptake of mineralized nitrogen. ¹⁴C-labelling has been used in studies of hydroponic (soil-free, nutrient-water) systems⁴, but not in the field. Double-labelled (¹³C, ¹⁵N) glycine was used in a macrocosm study of arctic tundra, showing that plant uptake of glycine-derived nitrogen was as fast as uptake of NH₄⁺, but this gave no evidence that the ¹³C-labelled tracer was taken up¹⁰. Variations in the natural abundance of ¹⁵N among plants in the field may be related to use of different nitrogen sources, but analysis of these variations is unlikely to yield conclusive quantitative evidence of the relative importance of organic and inorganic nitrogen uptake¹¹.

We aimed to test whether uptake of organic nitrogen occurs in the field, and to test whether plants with different types of mycorrhiza differ in this respect. We performed a low-tracer-level study with minimal disturbance of the soil-plant system in a late successional coniferous forest in northern Sweden (Table 1). Double, universally labelled (U) amino acid (U-13C2, 15N-glycine) injected at 1 kg nitrogen per hectare into the mor layer was used to study the importance of direct uptake of intact amino acid compared with uptake of nitrogen from mineralized amino acid, that is ¹⁵N-NH₄⁺. Typical European boreal forest species¹² (the ectomycorrhizal trees Pinus sylvestris L. and Picea abies (L) Karst., the ericoid-mycorrhizal dwarf shrub Vaccinium myrtillus L. and the arbuscular-mycorrhizal grass Deschampsia flexuosa (L) Trin.) were included in the study. Sequential sampling of above- and below-ground plant parts and mor layer soil was conducted after 6 h (2 h for above-ground parts), 1 day and 7 days. Respiration of ¹³C-CO₂ was measured after 2 h, 1 day and 7 days.

| Table 1 Characteristics of the mor layer at the site studied, Renberget | | | | |
|---|----------------|--|--|--|
| n∐* | 21+00 | | | |
| hu. | 3.1 ± 0.0 | | | |
| C:N ratio | 36.6 ± 0.9 | | | |
| Total soil nitrogen (kg ha ⁻¹) | 302 ± 33 | | | |
| Microbial nitrogen† (kg ha ⁻¹) | 27 ± 3 | | | |
| Extractable amino-acid nitrogen‡ (kg ha ⁻¹) | 1.0 ± 0.1 | | | |
| Extractable inorganic nitrogen‡ (kg ha ⁻¹) | 0.3 ± 0.05 | | | |

* 0.04 M CaCl₂, soil:solution ratio 1:10

† Method of ref. 20, using 0.45 as an estimate of the fraction of nitrogen released by fumigation. ‡ In 2 M KCl, soil:solution ratio 1:10.

Data (averages ± 1 s.e.m.) are from control plots (n = 22-24).

The results clearly show that a significant proportion of the supplied tracer was absorbed as intact amino acid. This is illustrated by plots of excess ¹³C against excess ¹⁵N in the soluble nitrogen fraction in roots (Fig. 1), and the highly significant regressions produced (P < 0.001). By comparing the slopes of these lines with the slope of 2, derived from the ¹³C:¹⁵N ratio of the tracer, a conservative estimate of the fraction of nitrogen taken up as amino acid can be obtained. After 6 h, this fraction was at least 42%, 64% and 91% for the trees, the grass and the dwarf shrub, respectively. Corresponding values after 1 day were 45%, 50% and 61%. Seven days after tracer application, there was no clear linkage between ¹³C and ¹⁵N levels in the trees and the grass, although the association was still seen in the dwarf shrub.

The slopes of the regression lines from the first samples show that the dwarf shrub had higher ¹³C:¹⁵N values than the trees (analysis of covariance followed by Tukey's test; P < 0.01) (ref. 13). Tracer ¹⁵N was also detected in shoots of the grass after only 2 h and in dwarf shrub shoots after 1 day (Table 2). The ¹⁵N-tracer level of both roots (Fig. 1) and shoots (Table 2) was similar for glycine-treated and NH₄⁺-treated plants, giving further support to the hypothesis that glycine-derived nitrogen was, to a large extent, taken up as the intact amino acid¹⁰.

Metabolism of glycine in plants or microorganisms occurs mainly by the route of serine synthesis, whereby 2 mol glycine are converted into 1 mol serine, 1 mol NH3 and 1 mol CO2 (ref. 14). Ammonium is re-assimilated, whereas CO₂ is presumably not re-assimilated in non-photosynthetic tissue. We used glycine labelled at both carbon

10 ECM tree 0 -5 ò 5 10 15 20 25 10 AM gras ³C excess (nmol per sample) 5 C -5 ò ò 10 EM dwarf shrub 5 0 Ġ

¹⁵N excess (nmol per sample)

positions, and the likely loss of the carboxy group during incorporation could explain the very rapid shift in the ¹³C:¹⁵N ratio from 2:1 to, or towards, 1:1 (Figs 1 and 2). Hence, our estimates of organic nitrogen uptake are probably underestimates, and differences between species could be partly related to differences in rates of decarboxylation. In the previous study of tundra plants¹⁰, only the carboxy carbon of the glycine used was labelled, which may explain the difficulties in detecting the ¹³C-labelled tracer. Here, some of the variability in natural abundance of ¹³C relates to differences between plants in shaded and unshaded environments¹⁵. Analysis of the chitin content¹⁶ of the roots indicated that all plants were mycorrhizal, with values of 0.5, 1.1, and 2.2 mg chitin per g root dry weight in the dwarf shrub, the grass and the trees, respectively. These data reflect the large proportion of rhizomes found in the samples of dwarf shrub roots. We cannot deduce how much of the ¹³C and ¹⁵N label was located in either plant or fungal cells. However, as the fungus is an integral part of the mycorrhizal root, this distinction is not critical in the context of biogeochemical nitrogen cycling. The most important consideration is whether or not organic nitrogen is taken up from the soil by the mycorrhizal roots.

Measurements in sifted soil from the labelled plots (Fig. 2) show that 64 \pm 2%, that is, most of the nitrogen applied as either glycine or NH₄⁺, was retained in this compartment. Unlike in previous studies, we did not prevent leaching and lateral flow or transport of glycine or NH₄⁺ by roots out of the labelled plots, to minimize disturbance to the studied system. The average ¹³C:¹⁵N ratio for the

| Table 2 | ∂ ¹⁵ N | of | shoots | from | the | dwarf | shrub | V. | myrtillus | and | the | grass |
|----------|-------------------|----|--------|------|-----|-------|-------|----|-----------|-----|-----|-------|
| D. flexu | osa | | | | | | | | | | | |

| Species | Treatment | | $\partial^{15}N$ of shoots* | | | |
|-------------|--|--|---|---|--|--|
| | | 2 h | 1 day | 7 days | | |
| Dwarf shrub | Control ¹⁵ NH ₄ (U- ¹³ C ₂ , ¹⁵ N) Glycine | $\begin{array}{c} -1.5\pm 0.6\\ 1.5\pm 1.5\\ -0.5\pm 0.7\end{array}$ | $\begin{array}{c} -1.5\pm 0.6\\ 2.5\pm 3.4\\ 3.3\pm 1.7\end{array}$ | - 1.1 ± 0.7 11.5 ± 7.0 3.8 ± 1.4 | | |
| Grass | Control ¹⁵ NH ⁺ (U- ¹³ C ₂ , ¹⁵ N) Glycine | $\begin{array}{c} - \ 0.1 \pm 0.9 \\ 2.1 \pm 0.8 \\ 2.6 \pm 1.4 \end{array}$ | -1.0 ± 0.7 12.5 ± 10 3.3 ± 1.5 | -0.3 ± 0.7 64.9 ± 36.5 20.5 ± 8.2 | | |

* Denotes part per thousand (‰) deviations form the ratio $^{15}\text{N}:^{14}\text{N}$ in atmospheric N_2. These results were obtained from plots treated with either water (control), $^{15}\text{NH}_4^*$ or (U- $^{13}\text{C}_2,$ ⁵N) glycine. (Averages ± s.e.m., n = 8.)





Figure 2 The relationship between excess ¹³C and excess ¹⁵N in the mor layer into which water, or solutions containing $(U^{-13}C_2, {}^{15}N)$ glycine or ${}^{15}NH_4^+$, was injected. The ¹³C:¹⁵N ratio of the glycine (2:1) injected (broken line) and the regression relating to glycine-injected plots are shown (unbroken line) (slope = 0.87, $r^2 = 0.67$). Data from all samplings are pooled as there were no major differences between samplings. Symbols as in Fig. 1; n = 24.

letters to nature

three samplings of the mor layer was 0.86, corresponding to 27% recovery of added glycine carbon (Fig. 2). The rapid decline in the ¹³C:¹⁵N ratio in the soil between 0 h and 1 day is corroborated by data showing that the ¹³C-CO₂ respiration rates were 14-fold higher at 2 h than at 7 days. These results further strengthen the conclusion that the turnover of glycine is very dynamic, and support the idea that uptake and initial decarboxylation of the glycine is rapid. The magnitude of that process (Fig. 2) indicates that biotic immobilization of glycine-derived nitrogen is at least as important as any possible mechanism of abiotic immobilization. For example, if all of the 64% of glycine recovered in the mor layer was decarboxylated, there should be 32% recovery of the added ¹³C. This is similar to the value found (27%, see above), which lends support to the idea that essentially all of the immobilization in the mor layer is biotic. In further support of this, release of ¹⁵N after fumigation and extraction, and retention of ¹⁵N-labelled tracer in the mor layer, were strongly correlated ($r^2 = 0.78$, P < 0.001). The biotic material of the sifted soil will include a fraction of ericaceous hair-roots and extramatrical mycelium, and therefore part of the nitrogen immobilized in this compartment is in the process of being transported to plants.

Evidence accumulated over a long period indicates that nitrogen is a key element regulating the production, structure and function of many terrestrial ecosystems¹. Our data indicate, however, that current conceptual models may be overstating the importance of nitrogen mineralization. We have shown that the dominant plant species in these forests, irrespective of their type of mycorrhiza, all compete well for, and use, glycine as a N source. This demonstrates that nitrogen mineralization is not necessary before plants use organic nitrogen, and suggests that a separation between these groups of plants into those that are capable and incapable of using organic nitrogen should be made with care. However, there may be quantitative differences applying to different organic nitrogen sources, in which the apparently exclusive capacity of ectomycorrhizal and ericoid-mycorrhizal plants to excrete proteases may play a key role⁶. Differences in the capacity of these species to respond to variations in nitrogen supply rate are probably also important¹⁷.

Methods

Field site. The experiment was conducted at Renberget (64° 14′ N, 19° 46′ E) in northern Sweden, 60 km northwest of Umeå. The area is located in the middle boreal zone in a late-successional *P. sylvestris/P. abies* forest of mesic dwarf shrub type, with a field layer dominated by the ericaceous dwarf shrub *V. myrtillus*. Other species present include the dwarf shrub *V. vitis-idaea* and the grass *D. flexuosa*. The total area used for the study was ~150 m × 50 m. The daytime temperature of the mor layer was ~7–8 °C during the study.

Experimental design. Circular plots (0.0464 m^2) were chosen so that shoots of both *V. myrtillus* and *D. flexuosa* plants were present within each plot. The site was divided into eight blocks, and within each block each treatment was replicated three times, allowing three samplings with eight repetitions each (n = 8).

Treatments. At 10:00 on 16 June 1996, the experiment was started by applying the labelled substances to the test plots, and water to the control plots. 250 ml 1.2 mM (U- $^{13}C_2$, 98 atom percentage; ^{15}N , 96–99 atom percentage) glycine or 1.2 mM ^{15}N -NH₄Cl (98 atom percentage) or water was injected into the mor layer (using a 50-ml syringe fitted with a 4-sideport needle) a few cm above the mineral soil. We made no attempts to stop solutions flowing either down into the mineral soil or laterally in the mor layer.

Sampling. The first step in sampling was cutting and recovering the aboveground vegetation on the plots. Thereafter, below-ground ${}^{13}C-CO_2$ respiration was measured using a cuvette technique¹⁸. These measurements were taken within 10 min, on each of six plots simultaneously. The below-ground compartment was sampled by cutting out the treated mor layer soil core. Samples were put into plastic bags and transported on ice to the laboratory. Above-ground structures of *V. myrtillus* and *D. flexuosa* were sorted so that only fresh young leaves were taken from both species. Roots of *V. myrtillus*, the two tree species (roots of which were not separated by species) and *D. flexuosa* were carefully picked from the soil core, thoroughly rinsed under tap water to remove soil material, and immersed thrice in 0.5 mM CaCl₂ to remove tracer adsorbed on root surfaces. Then, to obtain soil samples, the remaining cores were sifted through separate sieves (5-mm mesh) for each treatment. Samples were harvested (and stored at $-22 \,^{\circ}$ C) \sim 6 h (2 h for above-ground vegetation), 1 day and 7 days after application of the treatment solutions.

Analyses. Two sets of plant samples were prepared for analysis of ¹³C and ¹⁵N. The first set of plant samples and root-free soil samples were dried (70 °C, 24 h) and milled in a ball mill to a fine powder. The second set of plant samples was prepared by extraction of fresh and frozen roots in 10 mM phosphate buffer (pH 8.0), using 0.5 ml to 200 mg (fresh mass) plant material to obtain the soluble nitrogen fraction. The first set was used directly for isotopic analysis. The solutions of the second set were evaporated under reduced pressure to a final volume of \sim 50 µl, then transferred to tin capsules with Chromosorb absorbent. The extracts contained $\sim 2-5 \text{ mg}$ carbon and 0.02-0.05 mgnitrogen. The solutions were then dried (70 °C) before closing the capsules. Mass-spectrometric analysis was performed on a CF-IRMS¹⁹ (continuous flow isotope ratio mass spectrometer). Microbial nitrogen levels were determined in moist, root-free soil samples after one day of chloroform fumigation and extraction^{20,21}. The atom percentage of nitrogen released from microorganisms was measured after 48 h diffusion²² of distilled nitrogen from Kjeldahl extracts onto acidified Al₂O₃ in tin capsules. The diffusion was done in 12-ml gas-tight vials. Corrections for background nitrogen were made using a calculated blank²³, by comparing diffused and non-diffused isotope standards treated as the original samples. Concentrations of amino acids and inorganic nitrogen species in soil were determined in soil extracts (Table 1) by HPLC.

Calculations. Data on nitrogen and carbon content, and atom percentage ${}^{15}N$ and atom percentage ${}^{13}C$ in excess of the natural abundances of ${}^{15}N$ and ${}^{13}C$ in control plots were used to calculate mol ${}^{15}N$ excess and mol ${}^{13}C$ excess for each sample from plots treated with glycine or NH_4^+ .

Received 9 October 1997; accepted 5 February 1998.

- 1. Tamm, C.-O. Nitrogen in Terrestrial Ecosystems (Springer, Berlin, 1991).
- Vitousek, P. M. & Howarth, R. W. Nitrogen limitation on land and in sea. How can it occur? Biogeochemistry 13, 87–115 (1991).
- Melin, E. & Nilsson, H. Transfer of labelled nitrogen from glutamic acid to pine seedlings through the mycelium of *Boletus variegatus* (Sw.) Fr. *Nature* 171, 134 (1953).
- Chapin, F. S. III, Moilanen, L. & Kielland, K. Preferential use of organic nitrogen for growth by a nonmycorrhizal arctic sedge. *Nature* 361, 150–153 (1993).
 Abuzinadah, R. A. & Read, D. I. The role of proteins in the nitrogen nutrition of ectomycorrhizal
- Aduzinadan, K. A. & Read, D. J. The role of proteins in the nitrogen nutrition of ectomycorri plants. III. New Phytol. 103, 507–514 (1986).
- 6. Read, D. J. Mycorrhizas in ecosystems. Experientia 47, 376-391 (1991).
- Kaye, J. P. & Hart, S. C. Competition for nitrogen between plants and soil microorganisms. *TREE* 12, 139–143 (1997).
- Stark, J. M. & Hart, S. C. High rates of nitrification and nitrate turnover in undisturbed coniferous forest. *Nature* 385, 61–64 (1997).
- Eviner, V. T. & Chapin, F. S. II. Plant-microbial interactions. Nature 385, 26–27 (1997).
 Schimel, I. P. & Chapin, F. S. III. Tundra plant uptake of amino acid and NH⁴ nitrogen in situ: plants
- Schimel, J. P. & Chapin, F. S. III. Tundra plant uptake of amino acid and NH₄ nitrogen *in situ*: plants compete well for amino acid N. *Ecology* 77, 2142–2147 (1996).
- Högberg, P. ¹⁵N natural abundance in soil-plant systems. New Phytol. 137, 179–203 (1997).
 Hultén, E. & Fries, M. Atlas of North European Vascular Plants North of the Tropic of Cancer Vols 1–III
- Huiten, E. & Fries, M. Attas of North European Vascuar Plants North of the Tropic of Cancer Vols 1–111 (Koeltz Scientific Books, Königstein, 1986).
 Zar, J. H. Biostatistical Analysis, 2nd edn (Prentice-Hall, Englewood Cliffs, 1984).
- Oliver, D. J. The glycine decarboxylase complex from plant mitochondria. Annu. Rev. Plant Physiol. Mol. Biol. 45, 323–337 (1994).
- Ehleringer, J. R., Field, C. B., Lin, Z. & Kuo, C. Y. Leaf carbon isotope and mineral composition in subtropical plants along an irradiance cline. *Oecologia* 70, 520–526 (1986).
- Ekblad, A. & Näsholm, T. Determination of chitin in fungi and mycorrhizal roots by an improved HPLC analysis of glucosamine. *Plant Soil* 178, 29–35 (1996).
- 17. Chapin, F. S. III. The mineral nutrition of wild plants. Annu. Rev. Ecol. Syst. 11, 23-260 (1980).
- Högberg, P. & Ekblad, A. Substrate-induced respiration measured *in situ* in a C₃-plant ecosystem using additions of C₄-sucrose. *Soil Biol. Biochem.* 28, 1131–1138 (1996).
- Barrie, A. & Lemley, M. Automated N-15/C-13 analysis of biological materials. *Int. Lab. Techniques* 19, 82–91 (1989).
- Brookes, P. C., Landman, A., Pruden, G. & Jenkinson, D. S. Chloroform fumigation and the release of soil nitrogen: a rapid direct extraction method to measure microbial biomass nitrogen in soil. *Soil Biol. Biochem.* 17, 837–842 (1985).
- Jenkinson, D. S. in Advances in Nitrogen Cycling in Agricultural Systems (ed. Wilson, J. R.) 368–386 (CAB International, Wallingford, 1988).
- Risgaard-Petersen, N. & Rysgaard, S. in *Methods in Applied Soil Microbiology and Biochemsitry* (eds Alef, K. & Nannipieri, P.) 287–295 (Academic, London, 1995).
- Stark, J. M. & Hart, S. C. Diffusion technique for preparing salt solutions, Kjeldahl digests, and persulfate digests for nitrogen-15 analysis. *Soil Sci. Soc. Am. J.* 60, 1846–1855 (1996).

Acknowledgements. We thank M. Zetherström and G. Nyberg for assistance in the field and laboratory, and H. Wallmark for conducting the isotope analyses. This work was supported by grants from SJFR (to T.N.) and from NFR, SJFR and CEC (project CANIF) (to P.H.).

Correspondence and requests for materials should be addressed to T.N. (e-mail: Torgny:Nasholm@genfys. slu.se)